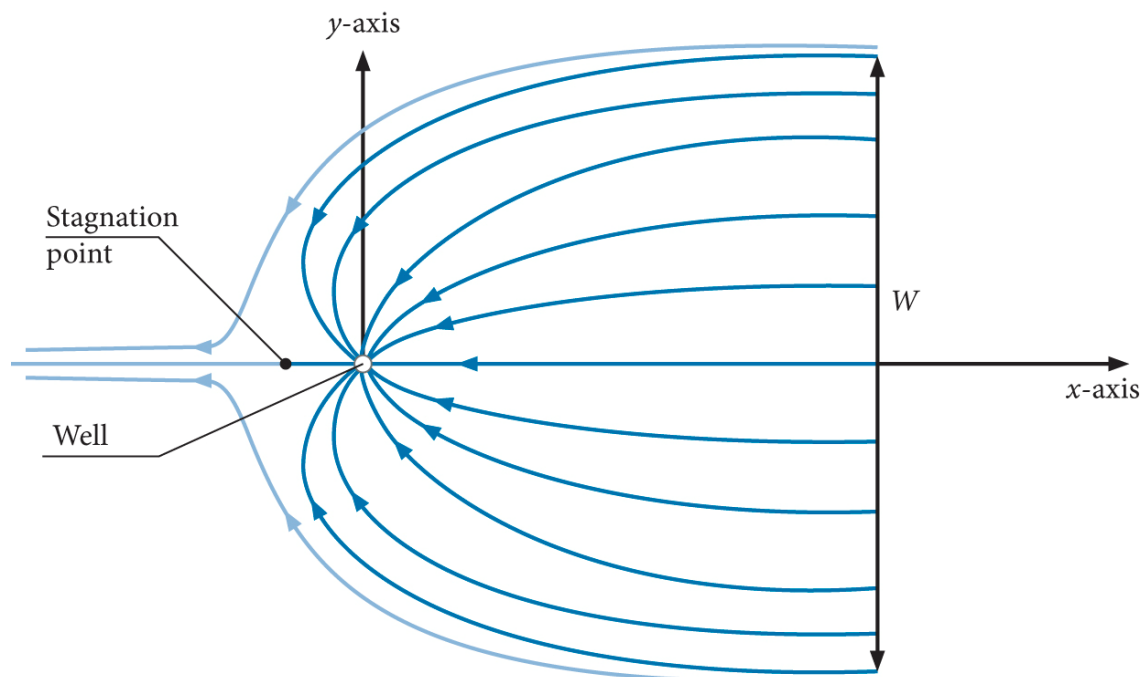


## Groundwater hydraulics test

A summary sheet of equations is provided on the last page of this exam; if possible, provide answers to this exam working your way from one or more of these equations; write down all your calculation steps!

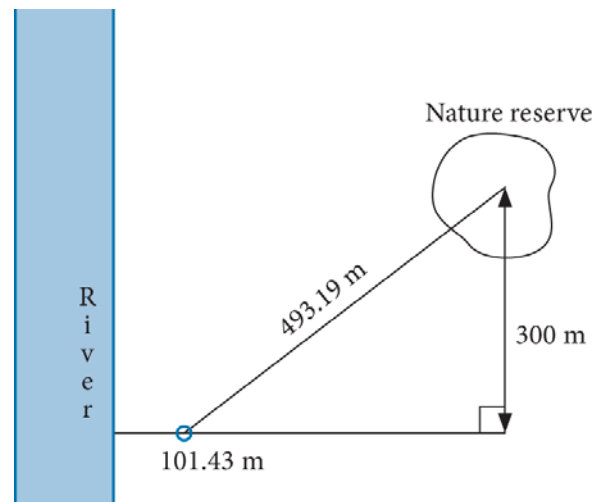
If you cannot solve a question, because you need the answer to another or earlier question, then choose a possible answer to this other/earlier question, clearly state on your exam paper what answer you have chosen and continue your calculations from there!

### Exercise 1 (2 credits)



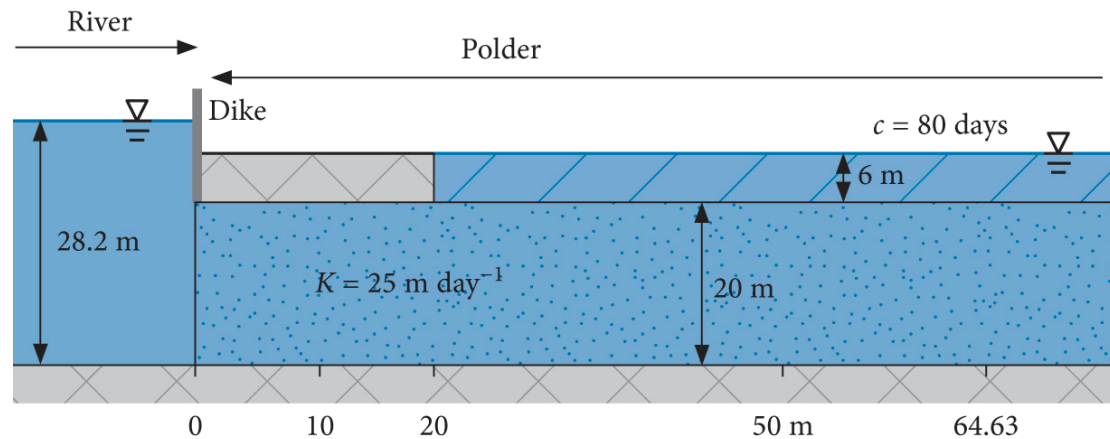
The above figure shows streamlines due to groundwater extraction by a pumping well in a regional flow field in plan view. As shown in the above figure a fully penetrating well (coordinates: 0,0) pumps water from a confined aquifer with a thickness of 50 m. Before pumping there existed a uniform groundwater flow parallel to the  $x$ -axis in a negative direction; the hydraulic gradient of this uniform flow field is 0.001. The aquifer has a hydraulic conductivity of  $10 \text{ m day}^{-1}$ . The well pumps water with a discharge or volume flux of  $628.32 \text{ m}^3 \text{ day}^{-1}$ .

- Determine the maximum width  $W$  of the area from which water is extracted by the well.
- Determine the location of the stagnation point in the groundwater flow field.

**Exercise 2** (2 credits)

On behalf of the drinking water supply, groundwater is pumped from a confined aquifer. The pumping discharge is  $314.16 \text{ m}^3 \text{ day}^{-1}$ . At a radial distance of 2000 m from the well there is no lowering of the hydraulic head due to pumping. The saturated depth of the aquifer is 50 m; the hydraulic conductivity of the aquifer equals  $10 \text{ m day}^{-1}$ . A river is located at a distance of 101.43 m of the pumping well, and the midpoint of a nature reserve is located at a distance of 493.19 m from the pumping well, both as shown in the above figure.

Determine the lowering of the hydraulic head in the midpoint of the nature reserve given the above setup.

**Exercise 3** (6 credits)

The above figure shows a typically Dutch inverse landscape profile ('Hollands profiel') with a high river, dike and low polder acting as a leaky confining layer. The water level of the fully penetrating river is situated 28.2 m above an impermeable layer. The aquifer overlies an impermeable layer, has a thickness of 20 m and a hydraulic conductivity of  $25 \text{ m day}^{-1}$ . The leaky confining layer has a thickness of 6 m and a hydraulic resistance to vertical groundwater flow  $c$  of 80 days. From  $x = 0$  to 20 m behind the dike the leaky confining layer has been made impermeable by sealing with clay. From  $x = 20$  m onwards the water table in the polder may be taken at surface level. All layers and aquifers are homogeneous and isotropic. Steady groundwater flow occurs in the  $x$  and vertical ( $z$ ) direction only.

Preferably use the storage function of your scientific calculator or else use sufficient figures/numbers after the decimal point in your calculations.

Determine the seepage into the polder between  $x = 20$  and 64.63 m in  $\text{m}^2 \text{ day}^{-1}$  using two methods, the 'horizontal-' as well as the 'vertical method'.

## Summary sheet groundwater hydraulics

### One-dimensional steady groundwater flow

Confined 
$$h = C_1x + C_2$$

Unconfined 
$$h^2 = C_1x + C_2$$

Leaky 
$$h = h_a + C_1e^{\frac{x}{\lambda}} + C_2e^{-\frac{x}{\lambda}} \text{ with } \lambda = \sqrt{K D c}$$

Recharge; equal water levels 
$$h^2 = -\frac{N}{K}x^2 + C$$

Recharge; different water levels 
$$h^2 = -\frac{N}{K}x^2 + C_1x + C_2$$

### Radial-symmetric steady groundwater flow

Confined 
$$h = h_R + \frac{Q_0}{2\pi K D} \ln \frac{r}{R} \text{ for } r_w \leq r \leq R$$

Unconfined 
$$h^2 = h_R^2 + \frac{Q_0}{\pi K} \ln \frac{r}{R} \text{ for } r_w \leq r \leq R$$